



ESRF

	Experiment title: Diffraction studied under extreme conditions of pressure and temperature	Experiment number: HS330
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Report:

During this beam-time we have performed 3 different experiments; (1) We extended the PVT equation of state (EOS) of MgSiO₃ perovskite up to 86 GPa and to 2700 K (2) We checked the occurrence of the orthorhombic polymorph of iron at high P and T, using either SiO₂ or Al₂O₃ as pressure transmitting medium, and Re or W as gasket material (3) We performed preliminary experiments on PVT-EOS of MgCl periclase, which can be considered as the second major constituent. of the Earth's lower mantle.

(PVT) equation of state of MgSiO₃ perovskite:

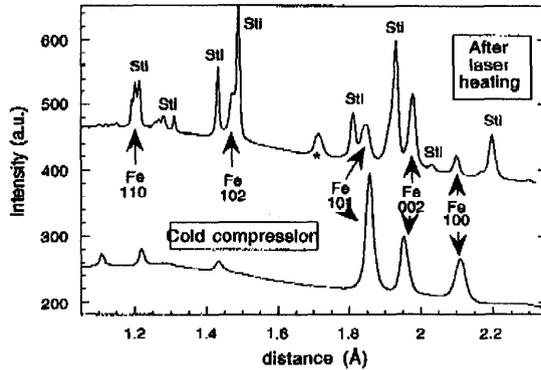
The experiments were carried out using a high power YAG multimode laser combined with a diamond-anvil cell. The use of this laser-type allows heating at higher pressure than the previously used CO₂ laser, but the experiments claim for a smaller X-ray beam, compatible in size with that of the hot spot (about 15 μm). We thus reached the extreme conditions of 86 GPa and 2700 K, that is about 2/3 of the maximum pressure in the Earth mantle. The thermoelastic parameters were determined after inversion of this PVT data set, and it represents to date the best experimental constraint to the mineralogical model of the Earth's lower mantle (Fiquet et al., 1998). Our goal was also to check the MgSiO₃ perovskite stability that was recently put into question. We found the perovskite stable, and it adopts the same Pbnm orthorhombic lattice at all P and T investigated in this study. The quality of the diffraction pattern even allowed us to carry out Rietveld full structure refinements (see experimental report HS325), which gave for the first time a very detailed structural behavior of this compound.

The orthorhombic structure of iron:

After our publication proposing an orthorhombic lattice for iron at high P and T [Andrault et al., 1997], we encountered large discussions about this structure solution. In particular, we received official comments that somehow asked for more proofs that our results were not related to experimental artifacts. These artifacts could have been due to (i) chemical reaction of our starting material, (ii) occurrence of Re gasket diffraction lines, (iii) deviatoric stress in the pressure chamber, and (iv) large temperature gradient.

We thus took time to reproduce the same orthorhombic lattice using different experimental conditions. We showed that the occurrence of the new characteristic orthorhombic diffraction lines is not due to any reaction with pressure transmitting medium or to gasket material, but to iron itself. We also showed that the temperature onto the sample is quite homogeneous with very reduced stress. This was done tracing the diffraction peaks FWHM as a function of pressure and temperature.

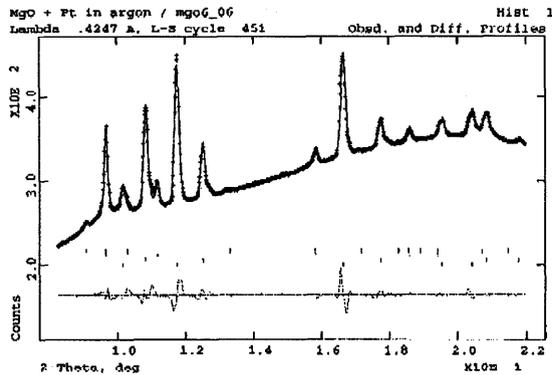
Figure 1: Integrated diffraction patterns of Fe in SiO₂ pressure medium, and using a W-gasket at 35 GPa. Upon cold compression (bottom spectrum), all e-hcp iron diffraction lines are visible, and there is no SiO₂-feature because quartz has become amorphous. The top spectrum was recorded after several laser-heating sequences. No oxidation of the iron-sample was encountered, since all diffraction peaks can be indexed as a mixture of iron and stishovite (* stands for metastable g-fcc iron, quenched on cooling).



Preliminary results on PVT-EOS of MgO:

There are several interests in the study of periclase at high P and T. It is the second major constituent of the Earth's lower mantle, and thus its PVT-EOS is needed to produce accurate mineralogical model for the deep Earth. MgO is often used as a reference in various high-pressure experiments, and thus one has to extend the P and T range in which its physical properties are well known. Finally, the mixing with FeO provides a useful solid solution to analyse the effect of iron on the elastic behavior of the MgO-rich minerals. X-ray diffraction patterns have thus been recorded in situ at high-pressure and high-temperature, on a mixture of MgO and platinum, loaded cryogenically in argon. Figure 2 shows an example of diffraction pattern, after collection and integration of an imaging plate exposed 120 s at a wavelength of 0.4247 Å.

Figure 2: In this experiment, the sample was compressed at 16.7 GPa while heated with a CO₂ infrared laser at around 1900 K, as determined by the analysis of the sample thermal emission. Such results, however, need to be extended to higher pressures and more investigations are needed to a slightly enriched iron composition, in order to identify correctly a reliable set of thermoelastic parameters



References:

Andraut et al., Science, 278, 831-834, October 1997
 Fiquet et al., Phys. Earth Planet. Int., in press, 1998